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Magnetic signal prospecting using multi parameter measurements

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Key words: Magnetic susceptibility, Magnetic viscosity, EMI, Magnetic signal, Mapping.

THE MAGNETIC SIGNAL OF SOILS

The magnetic signal of soils is divided between remanent (J_r) and induced magnetization (J_i). The former has diverse origins (heating, magnetic viscosity, slow deposition of magnetic bulk in a magnetic field) and is proof of the undisturbed state of features. The other, J_i , is acquired in the terrestrial magnetic field and is governed by magnetic susceptibility (κ). But the behavior of this property can be complex when the gain or loss of induced magnetization is decayed. This phenomenon corresponds to negative imaginary in complex κ ($\kappa = \kappa_{ph} - i\kappa_{qu}$), called magnetic viscosity (κ_{qu}).

The in-phase magnetic susceptibility, κ_{ph} , measurements are mainly sensitive to paramagnetic and ferrimagnetic materials. Beside detection of features, it can be used as a proxy for pedogenesis processes (heavy metal content or bacterial impact). κ_{qu} measurements are only sensitive to ferrimagnetic grains and particularly the near superparamagnetic (SPM)-single-domain (SD) ones.

From these primary measurements we can map derived parameters like the $\frac{\kappa_{qu}}{\kappa_{ph}}$ ratio which indicates the relative part of the magnetic viscous grain in the magnetic signal or the decrease slope map which indicates the spread of mag-

netic grain size distribution. With the κ_{ph} map, it is possible to evaluate the strength of induced magnetization (Benech *et al.*, 2002). Otherwise, from the κ_{qu} map, it is possible to evaluate the strength of the viscous part of remanent magnetization, assuming that grain size distribution is continuous and flat, the decrease slope is, then, near -1 (Thiesson *et al.*, 2007).

Fig. 1 summarizes the magnetic signal of soils and its parts. It also provides parameters that can be mapped. It must be noted that with few hypotheses and the measurements of κ_{ph} , κ_{qu} and the magnetic field anomaly, it is possible to differentiate J_r and J_i and to express the part of J_r resulting from viscous remanent magnetization of soil particles.

METHODS AND SURVEY

The Celtic site called 'les Arènes' located at Levroux (Indre, France) was occupied for about a hundred years between the middle of the second century BC and the middle of the first century BC. Its uniqueness lies in important craft activity and indeed archaeologists have found pits filled with metallurgical waste (Buchsenschutz *et al.*, 1988). We choose to survey an area called 'terrain Rogier' because of

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Figure 1: The magnetic signal of soil and how to measure it.

its archaeological and geological characteristics – pits for metallurgical waste were dug in a non-magnetic calcareous substratum, legitimizing therefore the use of both magnetic and electromagnetic methods. A G858 Cesium magnetometer with its two sensors at 0.4 m above the ground and spaced horizontally 0.6 m apart, records two parallel profiles simultaneously. The instrument was coupled with a CS60 EMI device (Job *et al.*, 1995). The data were recorded continuously along a profile, then transformed to fit a 1 x 1 m grid. The κ_{qu} maps were performed using VC100 (Thiesson *et al.*, 2007) with a 1 x 1 m grid.

RESULTS OF THE PROSPECTION

Different features can be recognized in the total magnetic field anomaly map (Fig. 2a): compact features with very strong anomaly magnitude and elongated ones with a lower anomaly magnitude. These two types of structures are also observed on the κ_{ph} map (Fig. 2b). On the κ_{qu} map (Fig. 2c), the compact features can be easily identified but the elongated ones do not appear clearly.

The magnetic anomaly map derived from the CS60 κ_{ph} , which corresponds to the top-soil induced magnetization anomaly, is clearly lower than the total magnetic field anomaly of the features. This can be the consequences of two facts: either the volume investigated by the CS60 does not contain the whole magnetic feature or the magnetic features have remanent magnetization. Nevertheless, based on the hypothesis that the superficial layer is the most disturbed and does not present any J_r , we subtracted the CS60 magnetic anomaly map from the total magnetic field anomaly one. The result is a magnetic anomaly map of undisturbed features.

From the value of the κ_{qu} it is possible to extrapolate the long term viscous remanent magnetization and to map the

Figure 2: Some maps of magnetic properties: a) Total magnetic field anomaly b) Magnetic susceptibility c) Magnetic viscosity at 44,1 μ s.

magnetic anomalies that are linked to long-term magnetic viscosity. Comparing the relative part of the viscous anomaly to the total magnetic anomaly ensures that we can distin-

guish features supposed to be linked to metallurgical activities (less viscous) and those more likely to be linked with settlements.

LABORATORY MEASUREMENTS

A set of samples was taken on both the features identified and a location free of features in order to compare the results from classical laboratory measurements and their interpretation with *in situ* data. The measurements consisted of:

- hysteresis cycles allowing assessment of grain size distribution;
- thermomagnetic measurements providing information on the nature of magnetic minerals;
- dual-frequency magnetic susceptibility measurements ensuring the evaluation of both κ_{ph} and κ_{qu} of samples.

The first laboratory results seem to show different features:

- samples taken on some strong magnetic anomalies show a very big magnetic susceptibility ($1300 \cdot 10^{-5} \text{SI} < \kappa < 3300 \cdot 10^{-5} \text{SI}$) and a very low frequency-dependence susceptibility with $\kappa_{fd}\%$ of about 1% at a depth of 0.4 m;
- samples taken on other strong circular magnetic anomalies show a big magnetic susceptibility of $500 \cdot 10^{-5} \text{SI} < \kappa < 1000 \cdot 10^{-5} \text{SI}$ but a high $\kappa_{fd}\%$ (between 8% and 10%) at a depth of 0.4 m. This indicates the presence of an important quantity of ferrimagnetic grains with sizes on the border between the SPM-SD ones (Dearing *et al.*, 1996);
- the superficial layer of soil and locations without features show lower susceptibility with $150 \cdot 10^{-5} \text{SI} < \kappa < 250 \cdot 10^{-5} \text{SI}$ and $3 < \kappa_{fd}\% < 8$.

Therefore, laboratory measurements seem to ensure a differentiation between metallurgical fillings showing small viscosity and other very viscous fillings which will be characterized by micromorphological analysis.

CONCLUSIONS

Three magnetic parameters (magnetic total field anomaly, magnetic susceptibility and magnetic viscosity) have been measured in this survey. These three parameters ensure a better characterization of the magnetic signal of soils. The simultaneous measurement of the total magnetic field and superficial magnetic susceptibility permits an evaluation of the part of the total magnetic field anomaly due to undisturbed features. The evaluation of the viscous part of the total magnetic field anomaly allows us to differentiate the type of the features (linked to metallurgy or to other activities). These results are reinforced by those obtained with classical laboratory measurements which confirm the reliability of the field results in terms of grain sizes distribution, for example.

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